Estimating and Correcting Global Weather Model Error



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Problem

- Inaccuracies in initial conditions and model deficiencies interact nonlinearly, causing numerical weather forecast errors to grow.
- With recent progress in data assimilation, the accuracy of initial conditions has improved dramatically.
- Accounting for model deficiencies has become relatively more important for data assimilation and ensemble forecasting.

Outline

- Brief review of model error correction
- SPEEDY model
- Generation of 6-hour forecasts and errors with NCEP reanalysis
- Separation of errors into monthly, diurnal, and state-dependent components
- Estimation and correction of model errors
- Results: our method is effective and computationally feasible
- Conclusions

Schemm et. al. (81, 86)

- introduced procedures for statistical correction of numerical predictions when verification data are only available at discrete times
- applying corrections only when verification data were available, they were successful in correcting artificial model errors
- reduced the small scale 12-hr errors of the NMC model
- errors at the larger scales grew due to randomization of the residual errors by the regression equations

Saha (1992)

- *nudged* a low-resolution version of the NMC operational forecast model to estimate systematic errors
- reduced systematic errors (measured against hi-res model) by adding artificial sources and sinks in heat, momentum, and mass
- correction *during* the integration and *a posteriori* correction were seen to give equivalent improvement in forecasts

Klinker and Sardeshmukh (1992)

- estimated average 6-hour forecast error from analysis increments of ECMWF operational model
- switched off each individual parameterization
- found that the model's gravity wave parameterization dominated the 1-day forecast error

D'Andrea and Vautard (2000)

- using a simple QG model, estimated tendency errors by solving for the model forcing which minimized 6-hour forecast errors
- during forecasts, found the closest analogues to the model state from the reference time series
- used the tendency errors corresponding to these analogues to correct forecasts online
- improved forecasts in Euro-Atlantic region
- computationally prohibitive

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$$\mathbf{g} = \dot{\mathbf{x}}_t - \mathbf{M}(\mathbf{x}_t) - \mathbf{L}\mathbf{x}_t - \mathbf{b}$$

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where \mathbf{x}_t is the state taken as truth (e.g. reanalysis)

• derived an empirical correction by minimizing $\overline{g^{\top}g}$ with respect to **b** and L

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DelSole and Hou (1999)

- applied Leith's procedure to a 2-layer QG model on an 8 x 10 grid (N=160 degrees of freedom)
- perturbed the model parameters to generate 'nature'
- resulting model errors were strongly state-dependent by design
- Leith's state-dependent error correction extended forecast skill
- computationally prohibitive for operational use

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IV. Correct the state-dependent errors.

SPEEDY Model, Molteni (2003)

- primitive equations, global spectral model
- contains parameterizations of condensation, convection, clouds, radiation, surface fluxes, and vertical diffusion
- T30 horizontal resolution, 7 sigma levels
- integrates vorticity, divergence, temperature, specific humidity, and surface pressure
- post-processed into horizontal wind, temperature, specific humidity, geopotential height, and surface pressure on 96x48 grid, 7 pressure levels
- dissipation and time-dependent forcing determined by climatological SST, surface moisture, albedo, land-surface vegetation, etc.

Generating Time Series of Model Forecasts and Errors





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Time Series and 5-year Climatology

- $\mathbf{x}_6^{f}(t)$ = time series of model states
- $\mathbf{x}_6^e(t)$ = corresponding 6-hour errors
- 5-year SPEEDY 6-hour climatology given by monthly mean $\mathbf{x}_6^{\mathrm{f}}$
- 5-year reanalysis climatology given by monthly mean $\overline{\mathbf{y}}$
- Bias given by monthly mean $\overline{\mathbf{x}_6^{e}}$

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200hPa Zonal Wind Monthly Bias

5-year Reanalysis Climatology $\overline{\mathbf{y}}$ (contour), Bias $\overline{\mathbf{x}_6^e}$ (color) January July



- SPEEDY underestimates zonal wind on the poleward side of the winter hemisphere jet.
- Exhibits large polar bias.

700hPa Specific Humidity Monthly Bias

5-year Reanalysis Climatology $\overline{\mathbf{y}}$ (contour), Bias $\overline{\mathbf{x}_6^e}$ (color) January July



• SPEEDY overestimates specific humidity at lower levels in the tropics, especially in the winter hemisphere over the oceans.

300hPa Specific Humidity Monthly Bias

5-year Reanalysis Climatology $\overline{\mathbf{y}}$ (contour), Bias $\overline{\mathbf{x}_6^e}$ (color) January July



• SPEEDY underestimates specific humidity at upper levels, especially in the summer hemisphere.

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Generate time series of 6-hour model forecasts and errors relative to the NCEP reanalysis using a simple but realistic GCM.

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2. Corrected <u>*a posteriori*</u>: Correct control forecast by 6-hour bias $\overline{\mathbf{x}_6^e}$ at 6 hours, 12-hour bias $\overline{\mathbf{x}_{12}^e}$ at 12 hours, etc.

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- 2. Corrected <u>a posteriori</u>: Correct control forecast by 6-hour bias $\overline{\mathbf{x}_6^e}$ at 6 hours, 12-hour bias $\overline{\mathbf{x}_{12}^e}$ at 12 hours, etc.
- 3. Corrected <u>online</u>: Integrate model, $\dot{\mathbf{x}} = \mathbf{M}(\mathbf{x}) + \frac{\overline{\mathbf{x}_6^e}}{\Delta t}$,

 $\overline{\mathbf{x}_6^{e}}$ is a daily linear interpolation (e.g. on July 1, $\overline{\mathbf{x}_6^{e}} = \frac{\overline{\mathbf{x}_6^{e}(Jun)} + \overline{\mathbf{x}_6^{e}(Jul)}}{2}$)

850hPa February 1987 Global Mean Anomaly Correlation



• Monthly bias correction gives substantial forecast improvement.

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• Monthly bias correction gives substantial forecast improvement.

• Online correction performs better than a posteriori correction.

Improvement of Online Correction Relative to Control



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• Improvements are uniform across levels in T, across seasons by level.

200hPa January 1982-86 Zonal Wind Bias Control Online Corrected Forecasts



• Most of the bias is removed by the online correction.

850hPa January 1982-86 Temperature Bias Control Online Corrected Forecasts



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Our Method: II. Diurnal Bias Correction

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- I. Estimate the monthly bias.
 - a. Compare the impact of correcting the model integration with statistical corrections performed *a posteriori*.
- II. Estimate and correct the diurnal errors.

- Anomalous model error time series: $\mathbf{x}_6^{e\prime}(i) = \mathbf{x}_6^e(i) \overline{\mathbf{x}_6^e}$
- Anomalous model error matrix: $D = [\mathbf{x}_6^{e\prime}(1) \ \mathbf{x}_6^{e\prime}(2) \ \dots \ \mathbf{x}_6^{e\prime}(N)]$
- The leading EOFs of DD[⊤] represent patterns of diurnal variability which are poorly represented by SPEEDY.

Leading EOFs of DD^{\top} , *T* at 925hPa, Jan 1982-1986



- Lack of diurnal forcing results in wavenumber 1 structure in the errors
- SPEEDY underestimates (overestimates) near surface daytime (nighttime) temperatures, more prominent over land

Principal Components

• Project leading EOFs onto anomalous errors (January, 1982)



- Leading pair of EOFs out of phase by 12 hours
- Find average strength of daily cycle over Jan 1982-86
- Compute diurnal correction as a function of the time of day

EOFs of DD^{\top}



• Diurnal correction substantially reduces error amplitude

Program of Applied Mathematics and Scientific Computation/University of Maryland

January 1982-1986

Diurnally Corrected 1987

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Our Method: III. State-Dependent Error Estimation

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Leith (1978) Empirical Correction Operator

- Anomalous analysis state time series: $\mathbf{y}'(t) = \mathbf{y}(t) \overline{\mathbf{y}}$
- Anomalous 6-hour error time series: $\mathbf{x}_6^{e\prime}(t) = \mathbf{x}_6^e(t) \overline{\mathbf{x}_6^e}$

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- Analysis state covariance: $\mathbf{B}_{yy}(t) = \mathbf{y}'(t) \ \mathbf{y}'^{\top}(t)$
- *Lagged* cross covariance: $\mathbf{B}_{\mathbf{x}^{e}\mathbf{y}}(t) = \mathbf{x}_{6}^{e'}(t) \mathbf{y}^{\prime \top}(t-1)$

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Leith's correction operator, given by $L = \overline{B_{x^ey}} \overline{B_{yy}}^{-1}$, provides a state-dependent correction:

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Problem: Direct computation of Lx requires $O(N^3)$ floating point operations *every* time step!

Implementation

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- This operation would be prohibitive for operational forecast models where N $\approx O(10^7)$.

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- 400 modes required to explain 90% of variance in dense L
- 40 modes required to explain 90% of variance in sparse L

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Principal Components: project error and state anomalies onto EOFs

$$a_k(t) = \mathbf{x}^{e'}(t) \cdot \mathbf{u}_k$$
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Heterogeneous correlation maps:

$$\rho[\mathbf{x}^{e\prime}, \mathbf{b}_k] = \left(\frac{\sigma_k}{\sqrt{\overline{\mathbf{b}_k^2(t)}}}\right) \mathbf{u}_k$$

III. State-Dependent Error Estimation

1982-86 error (color) and state (contour) coupled signals



• signals exhibit clear correlation in local structure

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IV. Correct the state-dependent errors.

The control model:

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The state-independent *online* corrected model:

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Leith's state-dependent corrected model given by:

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Our low-dimensional state-dependent corrected model is given by:

$$\dot{\mathbf{x}} = \mathbf{M}(\mathbf{x}) + \left[\overline{\mathbf{x}_{6}^{e}} + \sum_{k=1}^{K} \tilde{\mathbf{u}}_{k} \left(\frac{\mathbf{\sigma}_{k}}{\sqrt{\overline{\mathbf{b}_{k}^{2}}}}\right) \mathbf{v}_{k} \cdot \mathbf{x}'\right] \frac{1}{\Delta t}$$

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$$\sum_{k=1}^{K} \mathbf{\tilde{u}}_{k} \left(\frac{\mathbf{\sigma}_{k}}{\sqrt{\overline{b}_{k}^{2}}} \right) \mathbf{v}_{k} \cdot \mathbf{x}'$$

is the best representation of the original 6-hour forecast error anomalies $\mathbf{x}^{e\prime}$ in terms of the current anomalous forecast state \mathbf{x}^{\prime} .





We measure the forecast improvement using Leith's (univariate) dense and sparse correction operators and our low-dimensional approximation.

	Dense Leith	Sparse Leith	Low-Dim
Flops per time step	$O(N_{gp}^3)$	$O(N_{gp}^2)$	$O(N_{gp})$
Global Improvement	-8% (-4hr)	2% (1hr)	4% (2hr)
N. American Improvement	-6% (-3hr)	4% (2hr)	6% (3hr)

Chart contains average January 1987 improvement over state-independent corrected forecasts. Correction is more effective in regions where the heterogeneous correlations ρ are large.

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 - should work easily with existing data assimilation and ensemble schemes

Future

- Implement with data assimilation and ensemble schemes
- Test implementation on NCEP operational model (?)
- Reduce jumps in reanalysis climatology due to changes in observing system

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