Estimating and Correcting Global Weather Model Error

Chris Danforth, Eugenia Kalnay, Takemasa Miyoshi University of Maryland November 1, 2005

Problem

- Inaccuracies in initial conditions and model deficiencies interact nonlinearly, causing numerical weather forecast errors to grow.
- With recent progress in data assimilation, the accuracy of initial conditions has improved dramatically.
- Accounting for model deficiencies has become relatively more important for data assimilation and ensemble forecasting.

Outline

- Brief review of model error correction
- SPEEDY model
- Generation of 6-hour forecasts and errors with NCEP reanalysis
- Separation of errors into monthly, diurnal, and state-dependent components
- Estimation and correction of model errors
- Results: our method is effective and computationally feasible
- Conclusions

Schemm et. al. (81, 86)

- introduced procedures for statistical correction of numerical predictions when verification data are only available at discrete times
- applying corrections only when verification data were available, they were successful in correcting artificial model errors
- reduced the small scale 12-hr errors of the NMC model
- errors at the larger scales grew due to randomization of the residual errors by the regression equations

Saha (1992)

- *nudged* a low-resolution version of the NMC operational forecast model to estimate systematic errors
- reduced systematic errors (measured against hi-res model) by adding artificial sources and sinks in heat, momentum, and mass
- correction *during* the integration and *a posteriori* correction were seen to give equivalent improvement in forecasts

Klinker and Sardeshmukh (1992)

- estimated average 6-hour forecast error from analysis increments of ECMWF operational model
- switched off each individual parameterization
- found that the model's gravity wave parameterization dominated the 1-day forecast error

D'Andrea and Vautard (2000)

- using a simple QG model, estimated tendency errors by solving for the model forcing which minimized 6-hour forecast errors
- during forecasts, found the closest analogues to the model state from the reference time series
- used the tendency errors corresponding to these analogues to correct forecasts online
- improved forecasts in Euro-Atlantic region
- computationally prohibitive

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where \mathbf{x}_t is the state taken as truth (e.g. reanalysis)

• derived an empirical correction by minimizing $g^{\top}g$ with respect to **b** and L

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DelSole and Hou (1999)

- applied Leith's procedure to a 2-layer QG model on an 8 x 10 grid (N=160 degrees of freedom)
- perturbed the model parameters to generate 'nature'
- resulting model errors were strongly state-dependent by design
- Leith's state-dependent error correction extended forecast skill
- computationally prohibitive for operational use

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IV. Correct the state-dependent errors.

SPEEDY Model, Molteni (2003)

- primitive equations, global spectral model
- contains parameterizations of condensation, convection, clouds, radiation, surface fluxes, and vertical diffusion
- T30 horizontal resolution, 7 sigma levels
- integrates vorticity, divergence, temperature, specific humidity, and surface pressure
- post-processed into horizontal wind, temperature, specific humidity, geopotential height, and surface pressure on 96x48 grid, 7 pressure levels
- dissipation and time-dependent forcing determined by climatological SST, surface moisture, albedo, land-surface vegetation, etc.

Generating Time Series of Model Forecasts and Errors

1982-1986 NCEP Reanalysis

Generating Time Series of Model Forecasts and Errors

Program of Applied Mathematics and Scientific Computation/University of Maryland

Time Series and 5-year Climatology

- \bullet x $_{\epsilon}^{\mathrm{f}}$ $f_6(t)$ = time series of model states
- \bullet $\mathbf{x}_{6}^{\text{e}}$ $e_6^e(t)$ = corresponding 6-hour errors
- 5-year SPEEDY 6-hour climatology given by monthly mean x_f^f 6
- 5-year reanalysis climatology given by monthly mean \bar{y}
- Bias given by monthly mean $\overline{x_6^e}$ 6

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200hPa Zonal Wind Monthly Bias

5-year Reanalysis Climatology \bar{y} (contour), Bias \bar{x}_6^e $\frac{e}{6}$ (color) **January** July

- SPEEDY underestimates zonal wind on the poleward side of the winter hemisphere jet.
- Exhibits large polar bias.

700hPa Specific Humidity Monthly Bias

5-year Reanalysis Climatology \bar{y} (contour), Bias \bar{x}_6^e $\frac{e}{6}$ (color) January July

• SPEEDY overestimates specific humidity at lower levels in the tropics, especially in the winter hemisphere over the oceans.

300hPa Specific Humidity Monthly Bias

5-year Reanalysis Climatology \bar{y} (contour), Bias \bar{x}_6^e $\frac{e}{6}$ (color) January July

• SPEEDY underestimates specific humidity at upper levels, especially in the summer hemisphere.

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- 3. Corrected *<u>online</u>*: Integrate model, $\dot{\mathbf{x}} = M(\mathbf{x}) + \frac{\overline{\mathbf{x}_6^e}}{M}$ 6 ∆t ,

 $\overline{\mathbf{x}}_6^e$ $\frac{\overline{e}}{6}$ is a daily linear interpolation (e.g. on July 1, $\overline{\mathbf{x}_6^e} = \frac{\overline{\mathbf{x}_6^e}}{2}$ $\frac{e}{6}$ (Jun) + $\overline{\mathbf{x}_6^e}$ $\frac{\text{e}}{6}(\text{Jul})$ $\frac{+x_6(y_{\text{u}1})}{2}$

850hPa February 1987 Global Mean Anomaly Correlation

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● Online correction performs better than a posteriori correction.

Improvement of Online Correction Relative to Control

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• Improvements are uniform across levels in T, across seasons by level.

200hPa January 1982-86 Zonal Wind Bias Control Online Corrected Forecasts

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850hPa January 1982-86 Temperature Bias Control Online Corrected Forecasts

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Our Method: II. Diurnal Bias Correction

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- II. Estimate and correct the diurnal errors.

- Anomalous model error time series: $x_6^{e'}$ ${}_{6}^{\text{e}}(i) = \mathbf{x}_{6}^{\text{e}}$ $\frac{\text{e}}{6}$ (i) — $\overline{\mathbf{x}_6^{\text{e}}}$ 6
- Anomalous model error matrix: $D = [x_6^{e'}]$ $_{6}^{\text{e}}(1)$ $\mathbf{x}_{6}^{\text{e}}$ $6^{e'}(2) \ldots x_6^{e'}$ $_6^{\rm e\prime}({\rm N})]$
- The leading EOFs of DD^T represent patterns of diurnal variability which are poorly represented by SPEEDY.

Leading EOFs of DD^{\top} , *T* at 925hPa, Jan 1982-1986

- Lack of diurnal forcing results in wavenumber 1 structure in the errors
- SPEEDY underestimates (overestimates) near surface daytime (nighttime) temperatures, more prominent over land

Principal Components

• Project leading EOFs onto anomalous errors (January, 1982)

- Leading pair of EOFs out of phase by 12 hours
- Find average strength of daily cycle over Jan 1982-86
- Compute diurnal correction as a function of the time of day

$EOFs$ of DD ^T

Diurnally **Corrected** 1987

January

1982-1986

● Diurnal correction substantially reduces error amplitude

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Our Method: III. State-Dependent Error Estimation

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Leith (1978) Empirical Correction Operator

- Anomalous analysis state time series: $y'(t) = y(t) \overline{y}$
- Anomalous 6-hour error time series: $x_6^{e'}$ ${}_{6}^{\text{e}}(t) = \mathbf{x}_{6}^{\text{e}}$ $\frac{{\rm e}}{6}({\rm t})-\overline{{\bf x}_6^{\rm e}}$ 6

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- Analysis state covariance: $B_{yy}(t) = y'(t) y'^{\top}(t)$
- *Lagged* cross covariance: $B_{x^e y}(t) = x_6^{e}$ $_{6}^{\text{e}}(t)$ $\mathbf{y}'^{\top}(t-1)$

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Leith's correction operator, given by $L = B_{x^{e}y} B_{yy}$ −1 , provides a state-dependent correction:

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Problem: Direct computation of Lx requires $O(N^3)$ floating point operations *every* time step!

Implementation

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- This operation would be prohibitive for operational forecast models where $N \approx O(10^7)$.

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Sparse approximation of the Leith correction operator

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Explained variance of the SVD corresponding to *u* at 200hPa for the dense and sparse Leith operators.

- 400 modes required to explain 90% of variance in dense L
- 40 modes required to explain 90% of variance in sparse L

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Low-Dimensional Approximation based on regression

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Principal Components: project error and state anomalies onto EOFs

$$
a_k(t) = \mathbf{x}^{e\prime}(t) \cdot \mathbf{u}_k
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Heterogeneous correlation maps:

$$
\rho[\textbf{x}^{e\prime},b_k]=\Big(\frac{\sigma_k}{\sqrt{b_k^2(t)}}\Big)\textbf{u}_k
$$

III. State-Dependent Error Estimation

1982-86 error (color) and state (contour) coupled signals

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The state-independent *online* corrected model:

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Leith's state-dependent corrected model given by:

$$
\dot{\mathbf{x}} = M(\mathbf{x}) + \frac{\overline{\mathbf{x}_6^e}}{\Delta t} + L\mathbf{x}'
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Our low-dimensional state-dependent corrected model is given by:

$$
\dot{\textbf{x}}=M(\textbf{x})+\Big[\overline{\textbf{x}_{6}^{e}}+\sum_{k=1}^{K}\boldsymbol{\tilde{u}}_{k}\Big(\frac{\sigma_{k}}{\sqrt{b_{k}^{2}}}\Big)\textbf{v}_{k}\cdot\textbf{x}'\Big]\frac{1}{\Delta t}
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\sum_{k=1}^K \boldsymbol{\tilde{u}}_k \bigg(\frac{\sigma_k}{\sqrt{b_k^2}}\bigg) \boldsymbol{v}_k \cdot \boldsymbol{x}'
$$

is the best representation of the original 6-hour forecast error anomalies \mathbf{x}^{e} in terms of the current anomalous forecast state \mathbf{x}' .

We measure the forecast improvement using Leith's (univariate) dense and sparse correction operators and our low-dimensional approximation.

Chart contains average January 1987 improvement over state-independent corrected forecasts. Correction is more effective in regions where the heterogeneous correlations ρ are large.

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	- should work easily with existing data assimilation and ensemble schemes

Future

- Implement with data assimilation and ensemble schemes
- Test implementation on NCEP operational model (?)
- Reduce jumps in reanalysis climatology due to changes in observing system

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